#### Joint IMD-WMO group fellowship Training On Numerical Weather Prediction By Meteorological Training Institute, India Meteorological Department (IMD), Pune

# Atmospheric Boundary Layer (ABL) and its Parameterizations - Part III

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19 & 20 October 2021

# **1. References**

1. Turbulence in the Atmosphere – J. C. Wyngaard (Cambridge University Press) \*

2. An Introduction to Boundary Layer Meteorology - R. B. Stull (Kluwer Academic Publishers) \*

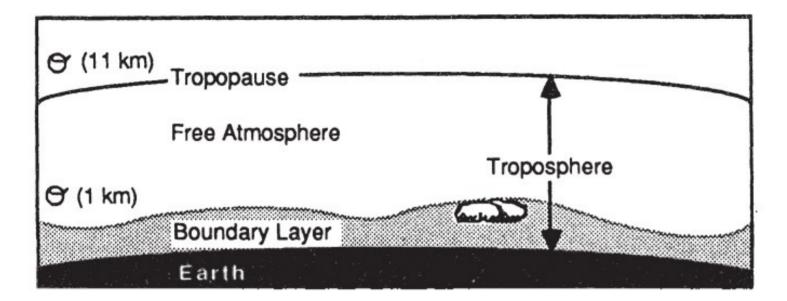
**3. An Introduction to Fluid Dynamics - G. K. Batchelor** (Cambridge University Press)

# **1.** References

4. Boundary-Layer Theory (7<sup>th</sup> Edition) – H. Schlichting (McGraw-Hill Book Company)

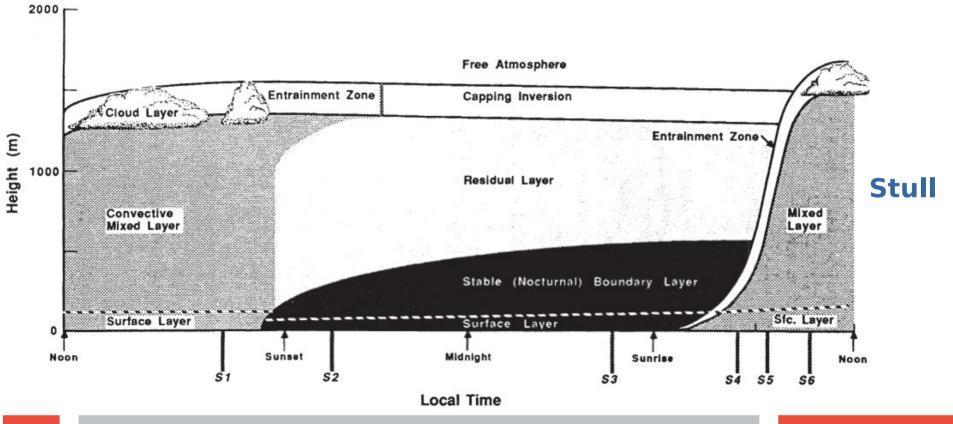
5. Turbulence - P. A. Davidson (Oxford University Press)

6. Fluid Mechanics - Kundu and Cohen (Academic Press)



Stull

Fig. 1.1 The troposphere can be divided into two parts: a boundary layer (shaded) near the surface and the free atmosphere above it.



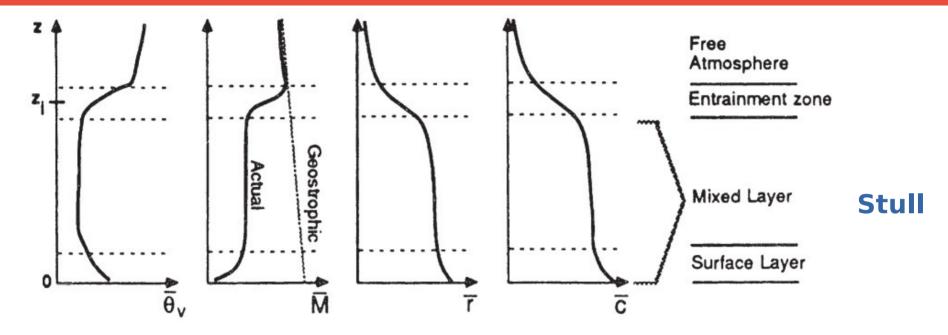


Fig. 1.9 Typical daytime profiles of mean virtual potential temperature  $\overline{\theta}_v$ , wind speed M (where  $M^2 = \overline{u}^2 + \overline{v}^2$ ), water vapor mixing ratio  $\overline{r}$ , and pollutant concentration  $\overline{c}$ .

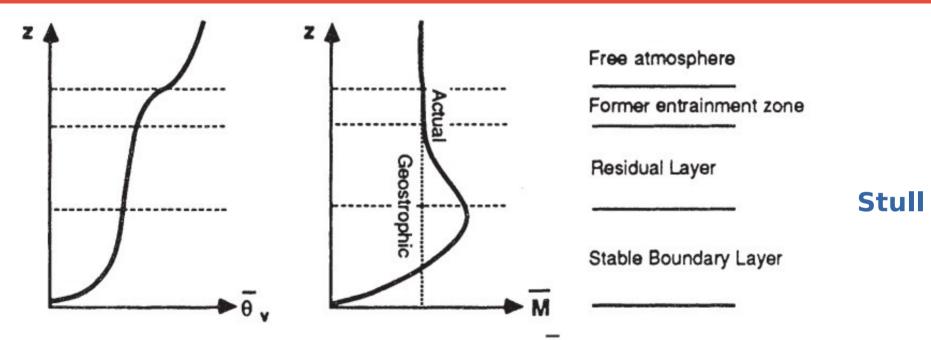
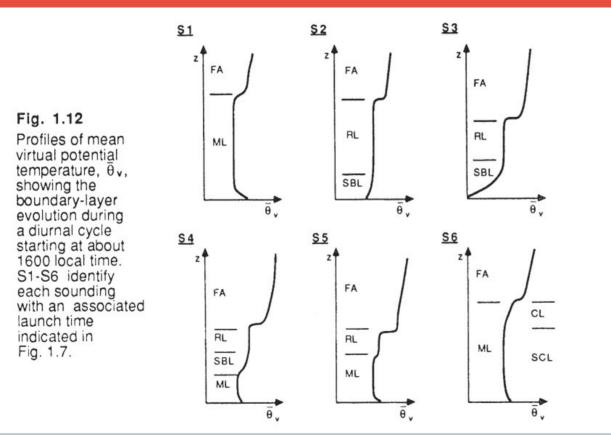
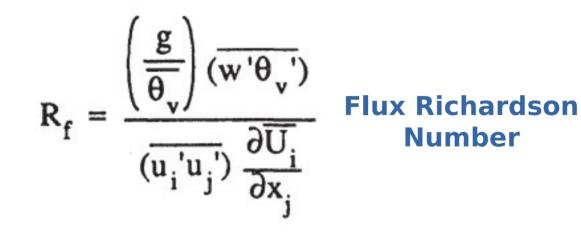
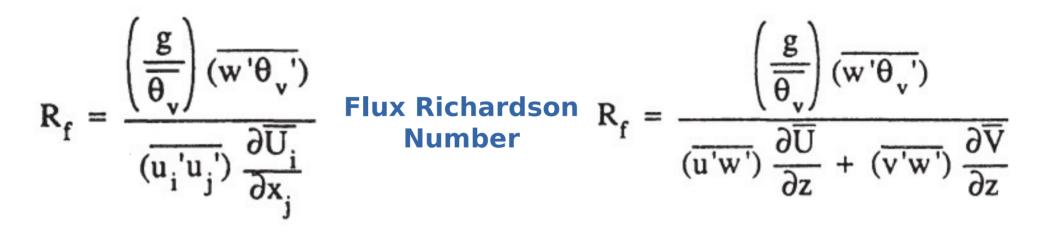


Fig. 1.11 Mean virtual potential temperature,  $\overline{\theta}_{v}$ , and wind speed,  $\overline{M}$ , profiles for an idealized stable boundary layer in a high-pressure region.

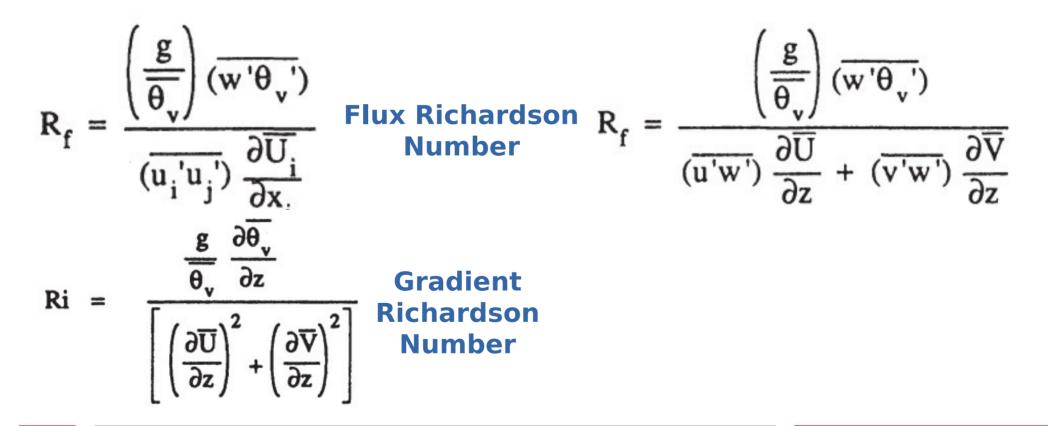


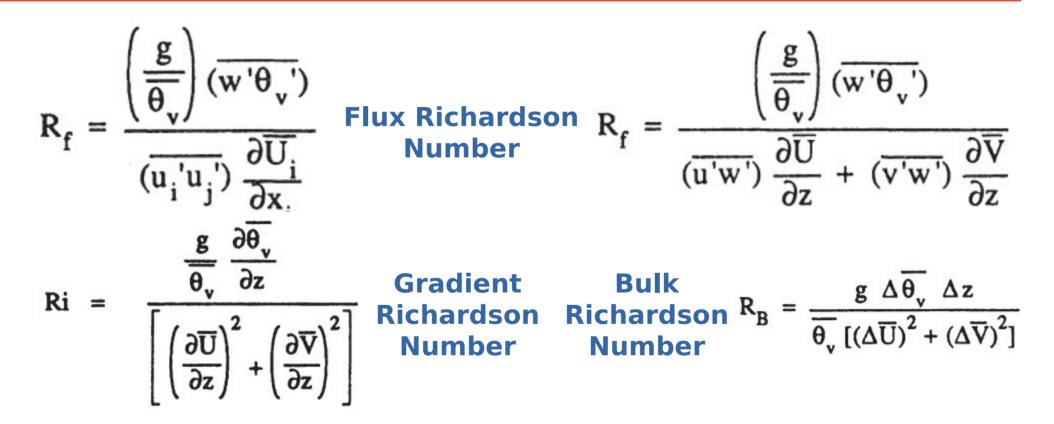
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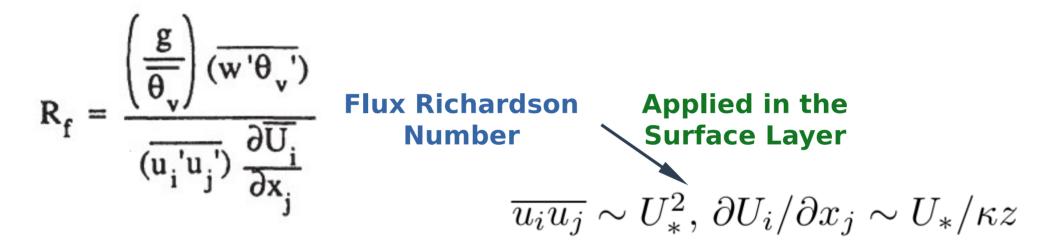


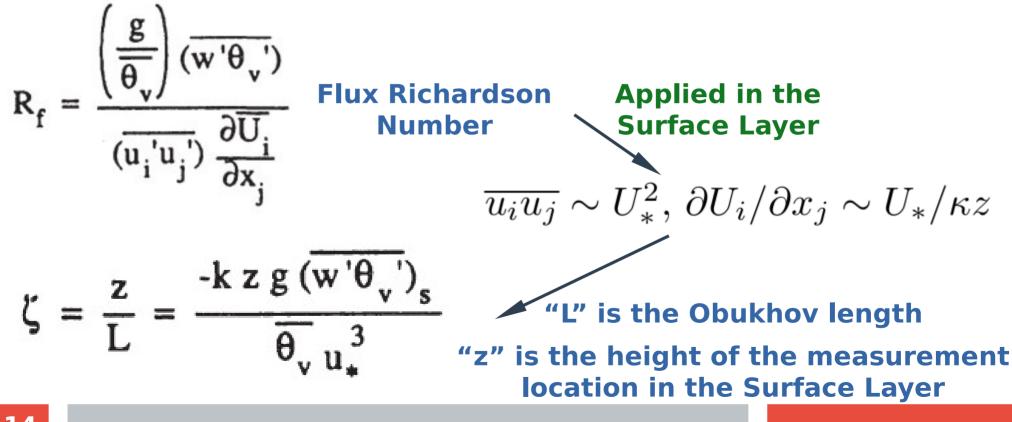
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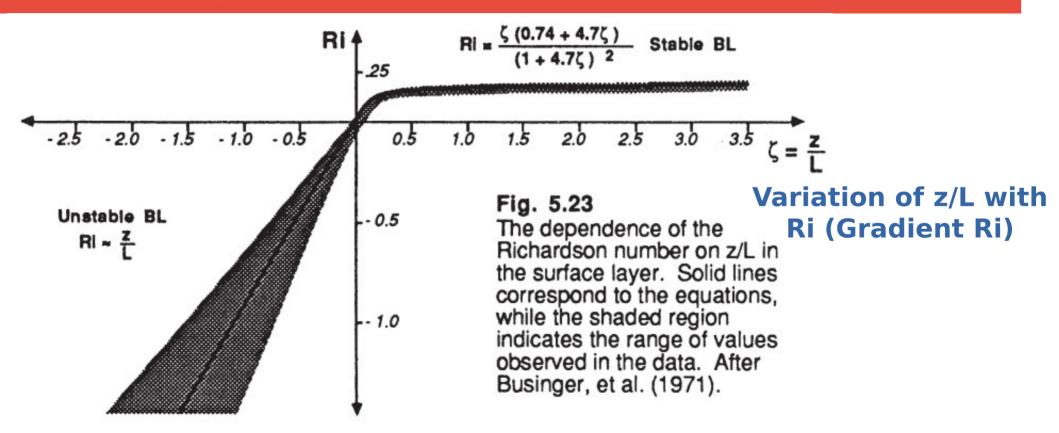
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$$L = -\frac{\overline{\theta_v} U_*^3}{\kappa g \left( \overline{w' \theta_v'} \right)_s} \qquad \qquad \zeta = \frac{z}{L} \qquad \qquad \zeta < 0, \text{ unstable} \\ \zeta > 0, \text{ stable}$$

"L" is the Obukhov length "z" is the measurement location in the Surface Layer Surface Layer stability parameter

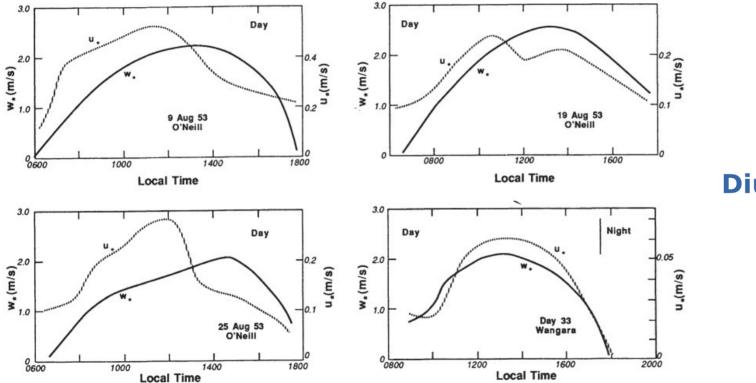


$$w_{*} = \left[\frac{g z_{i}}{\overline{\theta_{v}}} \left(\overline{w'\theta_{v'}}\right)_{s}\right]^{1/3}$$
Convective Velocity Scale  
zi is the height of the top of the ML

 $\overline{w'\theta'_v} > 0$ , convective i.e. heat flux upward from the surface

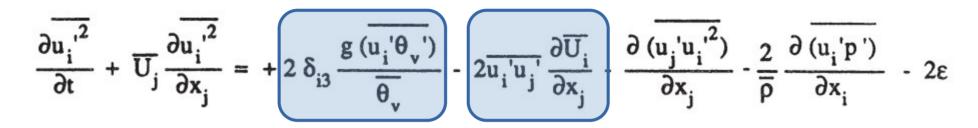
$$\zeta = \frac{z}{L} = -\frac{k z w_*^3}{z_i u_*^3}$$

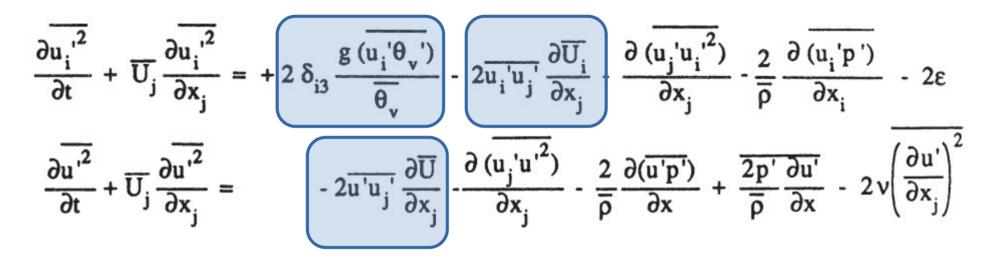
Surface Layer stability parameter in terms of convective and shear velocity scales

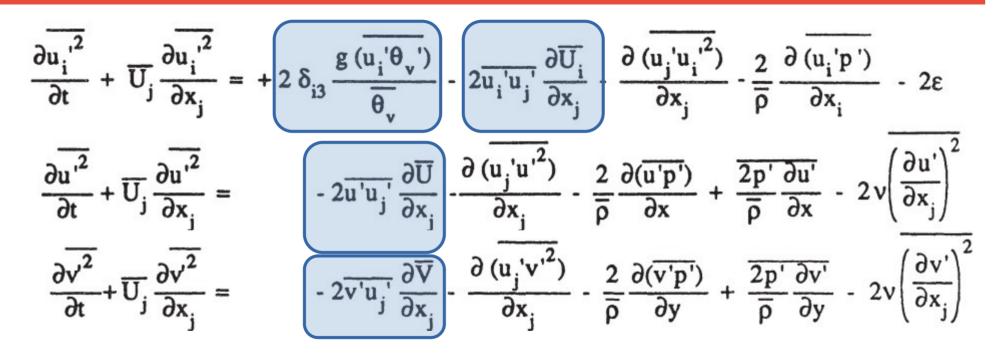


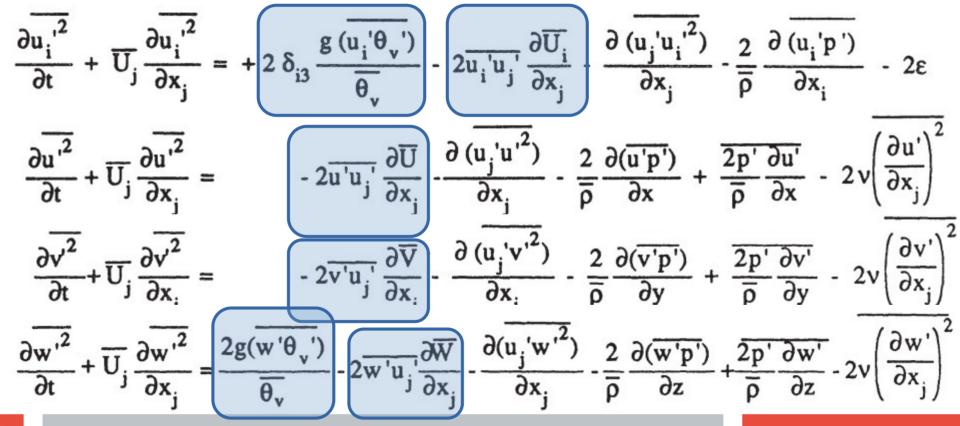
# Stull Diurnal variation of w\* and U\*

Fig. 4.1 Sample variations of the friction velocity, u<sub>\*</sub>, and the convective scaling velocity,w<sub>\*</sub>, with time for the O'Neill, (Nebraska) and Wangara (Australia) field programs.









#### What is the Surface Layer??

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- Close to a horizontally homogeneous surface,
- mean wind speed is small so that advection is negligible

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- Horizontal pressure gradient forces are weak, generally
- Therefore, the divergences of turbulent momentum and heat fluxes must be negligibly small OR the fluxes must be constant
- Surface Layer is the lowest few tens to hundreds of meter of the ABL where the turbulent momentum and heat fluxes are constant

Surface Layer closely corresponds to the Overlap Layer in Laboratory Turbulent Boundary Layers where the turbulent shear stress OR momentum flux is constant with respect to *z* 

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#### **Overlap Layer**

 $\frac{\partial U}{\partial z} = \frac{U_*}{\kappa z}$  $\frac{\kappa z}{U_*} \frac{\partial U}{\partial z} = \phi_M = 1$ 



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# **Overlap Layer**

#### Surface Layer

 $\phi_M$  is dimensionless wind shear

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# Overlap LayerSurface Layer $\frac{\partial U}{\partial z} = \frac{U_*}{\kappa z}$ $\phi_M$ is dimensionless wind shear $\frac{\kappa z}{U_*} \frac{\partial U}{\partial z} = \phi_M = 1$ $\frac{\kappa z}{U_*} \frac{\partial U}{\partial z} = \phi_M = f(\zeta)$ Proposal by Monin<br/>and Obukhov $f(\zeta = 0) = 1$ recovers the overlap layer scaling

# Dimensionless gradients in the Surface Layer

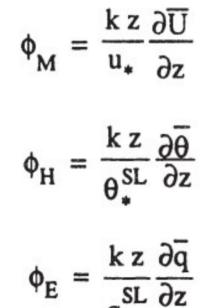
$$\phi_{\rm M} = \frac{k z}{u_*} \frac{\partial \overline{U}}{\partial z}$$

$$\phi_{\rm H} = \frac{k z}{\theta_*^{\rm SL}} \frac{\partial \overline{\theta}}{\partial z}$$

$$\phi_{\rm E} = \frac{k z}{q_*^{\rm SL}} \frac{\partial \overline{q}}{\partial z}$$

#### Dimensionless gradients in the Surface Layer Diff

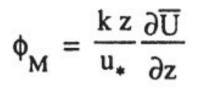
# **Different Classes of Similarity Theories**

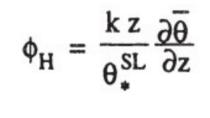


- Monin-Obukhov (Surface Layer) Similarity
- Mixed Layer Similarity
- Local Similarity
- Free Convection Similarity
- Rossby-Number Similarity

#### Dimensionless gradients in the Surface Layer Diff

# **Different Classes of Similarity Theories**



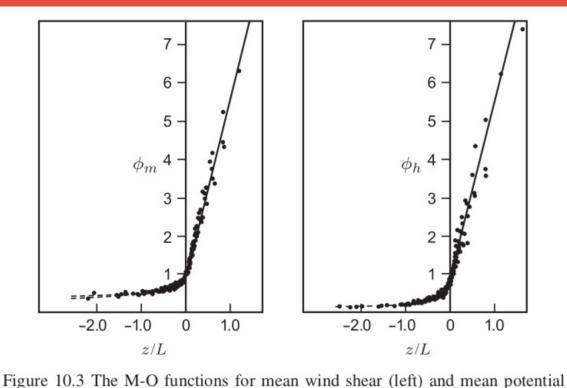


 $\phi_{\rm E} = \frac{\mathrm{k} \, \mathrm{z}}{\mathrm{sL}} \frac{\partial \overline{\mathrm{q}}}{\partial \mathrm{z}}$ 

Monin-Obukhov (Surface Layer) Similarity

- Mixed Layer Similarity
- Local Similarity
- Free Convection Similarity
- Rossby-Number Similarity

For Surface Layer Similarity  $\phi_M, \phi_H, \phi_E$  universal functions of  $\zeta$ 

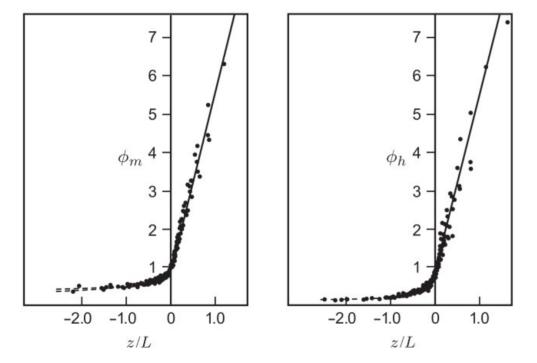


temperature gradient (right), Eq. (10.12), from the 1968 Kansas experiment. From

"Scaling" of Surface Layer data over a variety of stability conditions (Wyngaard)

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Businger et al. (1971).



"Scaling" of Surface Layer data over a variety of stability conditions (Wyngaard)

"Curve Fit" equations for different portions of these universal curves are the Surface Layer Parameterizations

(Stull, pg. 355+)

Figure 10.3 The M-O functions for mean wind shear (left) and mean potential temperature gradient (right), Eq. (10.12), from the 1968 Kansas experiment. From Businger *et al.* (1971).

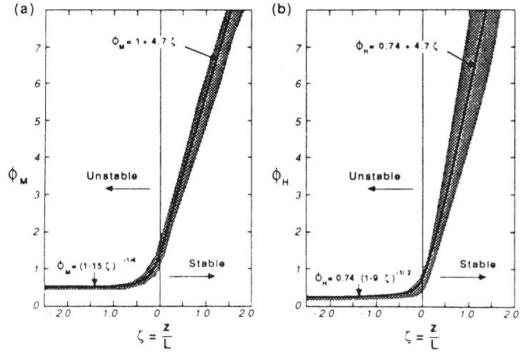


Fig. 9.9 (a) Range of dimensionless wind shear observations in the surface layer, plotted with interpolation formulas. (b) Range of dimensionless temperature gradient observations in the surface layer, plotted with interpolation formulas. After Businger, et al. (1971).

$$= 1 + \left(\frac{4.7 \text{ z}}{\text{L}}\right) \quad \text{for } \frac{\text{z}}{\text{L}} > 0 \quad (\text{stable})$$

$$\phi_{\text{M}} = 1 \quad \text{for } \frac{\text{z}}{\text{L}} = 0 \quad (\text{neutral})$$

$$= \left[1 - \left(\frac{15z}{\text{L}}\right)\right]^{-1/4} \text{ for } \frac{\text{z}}{\text{L}} < 0 \quad (\text{unstable})$$

$$= \frac{K_{\text{m}}}{K_{\text{H}}} + \frac{4.7 \text{ z}}{\text{L}} \quad \text{for } \frac{\text{z}}{\text{L}} > 0 \quad (\text{stable})$$

$$= \frac{K_{\text{m}}}{K_{\text{H}}} \quad \text{for } \frac{\text{z}}{\text{L}} = 0 \quad (\text{neutral})$$

 $\left. - \frac{9 z}{L} \right|^{-1/4}$ 

φ<sub>H</sub>

 $=\frac{K_{m}}{K_{H}}\left[1\right]$ 

for 
$$\frac{z}{L} < 0$$
 (unstable)

$$\tau_{viscous} = \nu \frac{\partial U}{\partial z}$$

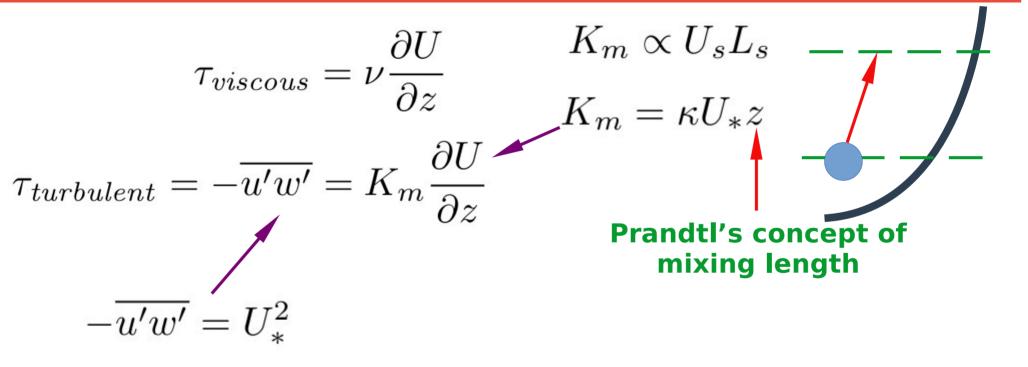
 $K_m \propto U_s L_s$ 

$$\tau_{viscous} = \nu \frac{\partial U}{\partial z}$$
$$\tau_{turbulent} = -\overline{u'w'} = K_m \frac{\partial U}{\partial z}$$

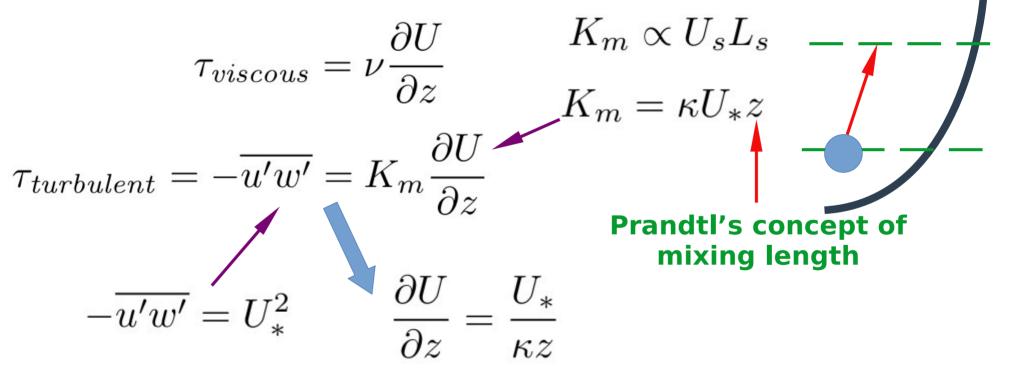
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$$K_m \propto U_s L_s$$
  
 $K_m = \kappa U_* z$   
Prandtl's concept of  
mixing length

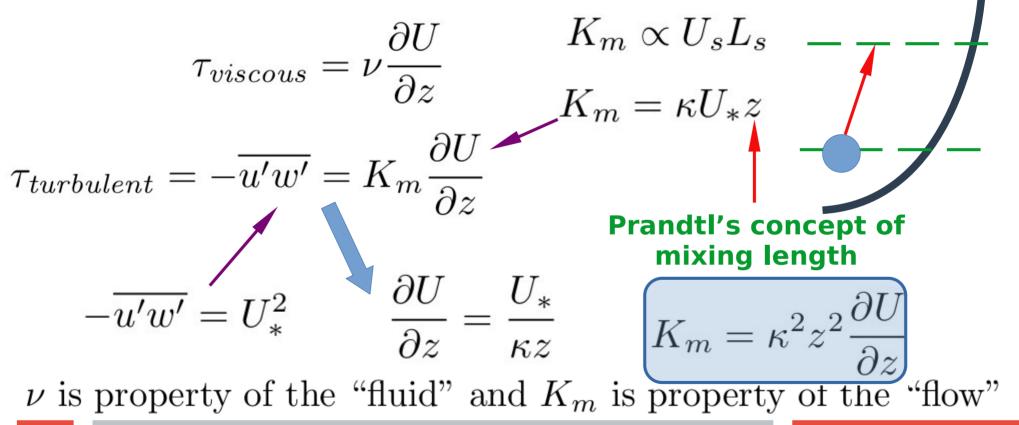
# $\nu$ is property of the "fluid" and $K_m$ is property of the "flow"



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# **6. Turbulence Closures or Parameterizations**

Page 197, Chapter 6 from Stull

